Towards a three-dimensional geological model of the North Sea subsurface

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ABSTRACT

The Geological Survey of the Netherlands is extending the 3D model of the Quaternary record as created for the onshore part of the country to the North Sea realm. Onshore, cores are the most important source of information. Offshore, seismic data are at least equally valuable. A recent pilot study has shown that 2D and 3D seismic data, originally collected for hydrocarbon exploration, are very useful in delineating much shallower stratigraphic units. This is best achieved by interpreting 3D seismic surveys together with high-resolution 2D seismic lines and well data. In the pilot, focus has been on the Oyster Grounds and Silver Well areas in the central part of the Dutch continental shelf. The units modeled thus far include Middle- to Late-Pleistocene strata that have accumulated in the North Sea Basin. The 3D seismic surveys reveal some remarkable, well-preserved sedimentary features that may be difficult to recognize in 2D profiles but are clearly visible in horizontal time slices. Examples include intricate early Holocene tidal-channel systems and striking sets of tunnel valleys formed during multiple glacial episodes. The level of detail provided by the time slices enables paleolandscape scholars to reconstruct environmental settings of Mesolithic (mostly Holocene) times.

INTRODUCTION

Geological models are quantitative, user-oriented predictions of subsurface architecture and properties in 3D space [Van der Meulen et al., in prep.]. They can be easily queried by end users of geological information to answer questions or make decisions in their respective areas of expertise. Two types of model have been developed for the onshore part of the Netherlands in recent years: (1) layer-based ones, with tops and bases of lithostratigraphic units only, and (2) voxel models composed of a regular grid of 3D pixels (voxels) with associated parameter values. These parameters include lithological characteristics, geotechnical, geochemical and hydrological attributes, and stratigraphy.

Geological modeling has replaced traditional mapping, which focused on producing qualitative visualizations of the subsurface. By using profile-type legends (‘Unit X on top of Unit Y on top of Unit Z’), maps depicting the shallow subsurface provide some information on the variability of sedimentary successions, but digital geological models are much better able to show this variability. Although the maps have considerable detail, application possibilities are limited because they are not tailor-made for specific purposes. They have been used mainly for illustrative purposes, and have contributed significantly to our understanding of process-response relationships and to the development of geological concepts, including aspects of preservation.

After a successful transition onshore, resulting in the models DGM, NL3D and GeoTOP developed by the Geological Survey of the Netherlands, offshore geological mapping is now being...
Table 1: Number and penetration depth of boreholes in pilot areas. Number of petroleum-exploration boreholes in parentheses.

<table>
<thead>
<tr>
<th>Location</th>
<th>Boreholes 0-5 m</th>
<th>Boreholes 5-12 m</th>
<th>Boreholes 12-50 m</th>
<th>Boreholes &gt;50 m</th>
</tr>
</thead>
<tbody>
<tr>
<td>Pilot area North</td>
<td>35</td>
<td>39</td>
<td>5</td>
<td>1 (11)</td>
</tr>
<tr>
<td>Pilot area South</td>
<td>788</td>
<td>282</td>
<td>111</td>
<td>3 (34)</td>
</tr>
</tbody>
</table>

replaced by digital geologic modeling as well. Protocols developed and lessons learned during the development of these onshore models are used as much as possible, but marine mapping and modeling offers its own specific challenges requiring adjustments in approach and methodology. These challenges are related to differences between onshore and offshore data types and data densities. The land data come primarily from boreholes, whereas the offshore has been mapped with a balanced combination of boreholes and geophysical surveys, even for the upper 50 m. Also, the density of boreholes (point data) decreases significantly with increasing distance from the shore (Table 1), forcing us to rely increasingly on 2D (cross-sections) and 3D (full coverage in x, y, z) seismic data. A second challenge associated with marine mapping is the dynamic nature of this environment. Much more than onshore, sediment is being transported to such an extent that maps of the shallow subsurface may be out of date after as little as a decade.

Here, we report on the development of a protocol for modeling the shallow (upper 100s of m) subsurface of the Dutch part of the North Sea. Efforts to optimize data density are highlighted, as they determine the scale at which lithostratigraphic and lithological units can be recognized and mapped in light of the limited number of boreholes. Using this new protocol, lithological horizons are being digitized to form a layer-based model for two pilot areas (Figure 1), and initial steps are undertaken to build a voxel model that will allow seamless integration with onshore voxel models.

**DATA TYPES**

**Borehole logs and laboratory analyses**

Borehole logs are extracted from the DINO database developed and hosted by the Geological Survey of the Netherlands. Most information comes from shallow boreholes, typically up to 12 m deep but locally up to several 100s of m. The logs of some hydrocarbon-exploration wells provide information for the upper 100s of m as well. Most boreholes up to 5 m deep provided little-disturbed core samples. The vast majority of deeper boreholes were made with counterflush or related techniques resulting in disturbed samples. Per depth interval, borehole logs provide qualitative, visually assessed information on lithology and stratigraphy. For many boreholes, one or more samples were analyzed in the laboratory, most frequently for quantitative measures of grain-size parameters, micropaleontology or age.

**Seismic profiles and time slices**

In marine seismic-reflection surveys, acoustic waves generated by a seismic source towed by or mounted to a vessel are used to obtain information on the subsurface. When a seismic wave encounters a boundary between two materials with different acoustic properties, some of the energy in the wave will be reflected at the boundary, while some of the energy will continue through the boundary. The amplitude of the reflected wave depends on the impedance contrast between the two materials, which is greatest at the water-seabed interface. A distinction can be made between high-resolution single- or multi-channel systems using frequencies of 100-1000 Hz and low-resolution multi-channel systems using <100 Hz. High-frequency single-channel systems record through short hydrophones. Their vertical resolution is on the order of a few meters, but they suffer from quality loss below first seabed multiple. High-frequency multi-channel systems record through hydrophone arrays in cables (streamers). Vertical resolution is on the order of meters below 100 ms (100 ms two-way travel time (TWTT) ≈ 75-80 m, including the water column), but lower in shallower units. Low-frequency multi-channel systems record through km-long hydrophone arrays. Vertical resolution is 5-15 m, depending on the frequency, on survey and source characteristics in relation to target geometry and on post-survey processing [Praeg, 2003]. In general, information of the upper 100-200 ms is lost. When acquisition parameters and processing are optimized for the shallow subsurface, however, information may be obtained from the seabed downward (starting at about 50 ms) when water depth is at least 40 m. In 3D surveys, multiple streamers are deployed in parallel to record data suitable for 3D interpretation of the structures beneath the seabed. Most 3D surveys employ low-frequency systems, which work best in deeper waters and cover very large areas efficiently at a coarse resolution [Fitch et al., 2005].

Data collected via a single cable or streamer are displayed as vertical cross-sections through the subsurface. In our traditional surveys conducted for mapping purposes, distances between adjacent profiles are typically on the order of 10 km or more. Such a wide spacing is commonly sufficient to understand the overall architecture and formation of sedimentary systems, and may be good enough for pre-planning purposes, but render maps unsuitable for most applied use. Not all reflections can be correlated among lines, and most architectural elements have lateral dimensions smaller than can be resolved in a 10 x 10 km grid.

Data collected with multiple parallel streamers may be displayed as vertical cross-sections, time slices or horizon slices. Time slices are sections of 3D seismic data having a certain arrival time. Because of spatial variability of sound velocity in the subsurface, they are near-horizontal rather than perfectly horizontal depth slices. Because of their map view, time slices are very suitable to identification of landscape elements as reflected in seismic amplitude. Horizon slices show the spatial pattern of particular reflections, created by tracing these reflections on all survey lines and interpolating the resulting data. Thus, features
can be extracted not only in planform but also in 3D, providing depth (and relative age) relationships.

**WORKFLOW**

**Stratigraphic framework**

The stratigraphic framework for the pilot studies is provided by published 1:250,000 map sheets made jointly by the British, Belgian and Dutch geological surveys during the 1980s and 1990s, and by the lithostratigraphic overview of Rijsdijk et al. [2005]. The borehole and seismic data are used to identify and map the tops of stratigraphic units. In identifying these units, well-defined reflections of 2D and 3D seismic datasets are linked and labeled systematically from the mid-Miocene unconformity upward, focusing on vertical (TWTT) mismatches at line intersections. These mismatches relate in part to positioning inaccuracies of older surveys.

**Petrel project**

At the core of the interpretation activities, a mother file was created in Petrel, a software application for the visualization and analysis of aggregated reservoir data from multiple sources. This mother file hosts all interpreted horizons. For interpretation, elements of the mother file are copied onto local hard drives. Changes and additions made on such local copies are exported regularly to the mother file.

Although the 3D seismic data allow superior visualization of even small sedimentary units, they are not ideally suited for the development of an overall stratigraphic framework of shallow (<300 m) units in Petrel. Therefore, we started reinterpreting high-quality 2D seismic lines collected during the 1980s and 1990s for the 1:250,000 maps. These lines were labeled, linking the major reflections to boundaries between stratigraphic units, and providing metadata including certainty of interpretation and name of interpreting geologist. The labeling scheme (for both bounding reflections and seismic facies) is summarized and updated in a table that includes information on reflection strength, internal reflection configuration, external form of bounded unit, degree of transparency and position relative to the reference datum (mean sea level).

The resulting framework is validated and supplemented with 3D seismics. By scrolling up and down in 3D seismic cubes at intersection planes with 2D seismic profiles, seismic facies and reflections identified on the 2D data can partly be assigned to depositional settings and environments. More subtle features on the 2D lines that were missed initially are labeled as part of this validation step. Finally, the 3D data are interpreted in their own right.

The seismic interpretations are validated with borehole data and with the results of laboratory analyses.

**Export to modeling software (ISATIS)**

The horizon slices as reconstructed in Petrel are exported as x-y-t grids with a time rather than a depth value for each x-y coordinate. The t values are translated into z values using seismic-wave velocities of x ms⁻¹. The resulting output is imported into Isatis, geostatistical software used to generate a digital geological model.

**WORKFLOW OUTPUT**

**Pilot area North**

The northern pilot area offers high-quality 3D seismics supplemented with 2D seismics and few boreholes. Using the workflow developed as part of the pilot study, we reconstructed the top of the Lower-Pleistocene Markams Hole Formation for an area of 55 x 65 km (Figure 2). When viewed at this large scale, this horizon slice provides a general impression of highs and lows, with an overall deepening from the southwest to the northeast. The eastern half of the image shows more complicated patterns associated with faults.

When zooming in and intersecting seismic profiles or horizon slices with time slices, sub-kilometer-scale patterns visible on the time slices provide direct indications of depositional environments. Some features can be easily recognized and interpreted, such as tunnel valleys, iceberg-scour marks and various channel fills. Tunnel valleys are large, valleys formed by meltwater under Pleistocene ice sheets near their margins. They may be up to 100 km long, kilometers wide and 100s of meters deep. The scour marks of icebergs, formed in open-marine waters beyond ice-sheet margins, appear as thin lines, commonly with a fairly uniform direction. Channel fills show intricate meandering and dendritic patterns (Figure 3). They may be tidal or fluvial in origin, and Pleistocene or Holocene in age.

The clear visibility of some of the smaller features on time slices is most likely related to the presence of shallow gas [cf. Schroot and Schüttenhelm, 2003], either formed in situ (e.g., organics in tidal-channel fills) or supplied from an external source.

![Figure 2. Horizon slice of Markams Hole Formation in pilot area South. Darker gray tones denote lows.](image-url)
Digital geological models that include 3D seismic data are superior in detail to those relying exclusively on 2D seismic data and boreholes. For most end users of geological information, such detail is essential. Marine archaeologists, for example, rely heavily on time-slice interpretation for paleolandscape reconstructions [Fitch et al., 2005]. They need to be able to recognize former coastlines, fertile lowlands, and drainage patterns and directions.

The availability of 3D seisms for large parts of the North Sea, related to the abundance and distribution of petroleum reservoirs, provides a unique opportunity to model the shallow subsurface at a spatial resolution comparable to that of the onshore voxel models NL3D (250 x 250 x 1 m) and GeoTOP (100 x 100 x 0.5 m) and layer-based model DGM (100 x 100 m). This similarity in resolution will make it easier to work toward an integrated land-sea model of the shallow subsurface. Such a model will help understand subsurface-related processes that are unrelated to the present North Sea shore. Hydrological models, for example, may be optimized when the distribution of units with various hydraulic properties is known both on- and offshore. The creation of a digital geological model for the entire North Sea, and its integration with existing or future onshore models for North Sea countries will be a key step in the development of a web-accessible geo-model for all of Europe.

CONCLUSIONS

The joint use of 2D and 3D seismic data is very useful in delineating shallow stratigraphic units, revealing features that may be difficult to recognize in 2D profiles. A new workflow is used to delineate unit boundaries and to map features as small as 10s of meters laterally and as thin as a decimeter. The presence of shallow gas enhances many of these features but also leads to vastly overestimated apparent thicknesses. The resulting improvement in mapping resolution is a key step toward an integrated land-sea model of the shallow subsurface.

REFERENCES

Fitch, S., K. Thomson, and V. Gaffney (2005), Late Pleistocene and Holocene depositional systems and the paleogeography of the Dogger Bank, North Sea, Quaternary Research, 64, 185-196.


